Detectability of subsurface oceans and their flows by JUICE

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Overview of magnetic induction at Galilean moons



including electrodynamics and ocean motion

Jovian B-field at Ganymede

- Main oscillation at the ~10h Jupiter rotation period
- Additional finer oscillations are expected:
 - 171h orbital period of Ganymede
 - 8h/15h period due to envelope behaviour of current sheet hinging
- Several additional harmonics at multiples of these frequencies





Variation of Jovian B field at Ganymede

- B field model from Connerney (1993) + Khurana (2005)
- 10h period oscillation (dipole tilt)
- small perturbations at orbital period due to orbit eccentricity and day/night differences
- B_Z along Jovian rotation axis
- \bullet B_{ρ} along Jupiter-Ganymede axis
- \bullet B_φ along Axis tangent to orbit
- Notice several frequencies of interest:
 - ω_{J} Jovian rotation frequency
 - ω_G Ganymede's orbital frequency
- red: linear combinations of these frequencies and their harmonics



Fourier analysis of the background jovian magnetic field at Ganymede

Induction response at Ganymede

B-field



E-field



Induction E-field at Europa and Ganymede

Europa





Ganymede



Bounds on long period oscillations

Ocean characteristics are chosen so as to obtain an 84% response at the Jovian rotation frequency (as measured by Kivelson)

Two extreme cases:

- Surface ocean with low conductivity (A)
- Deep and infinitely conductive ocean (B)

Finer frequencies will allow determining ocean depth and conductance



Inducing field strength	
ω _J	90 nT
2ω _J	5 nT
3ω _J	6 nT
ω _G	0.8 nT
2ω _G	0.5 nT
3ω _G	0.1 nT

Ocean magnetic induction on Earth



- Ocean water flows have been shown to induce a detectable B-field
- Is the same happening on Ganymede and Europa?
- This would allow us to remotely constrain ocean flows

MITgcm ocean model

- We have adopted MITgcm to simulate the oceans of Ganymede and Europa
- Ocean motion driven by gravitational tides and internal heat sources (equatorial heat flux on seafloor with maximum of ~10 mW/m²)



Temperature



Ocean tides (gravitational)

- Ocean tides are induced gravitationally
- These lead to radial and horizontal flow of the water
- This, in turn, will induce a B-field
- Ocean flow can also be generated by internal heat sources (not shown)
- We have adopted the MIT-GCM to Ganymede to calculate these flows and associated induced fields

Radial displacement



-150

-100

-50

0

Longitude (°)

50

100

150

Horizontal flow



B-field induced by ocean tides (gravitational)

 Ocean flows through Ganymede's fixed dipole induces a small magnetic perturbation b_r:

$$\left(i\Omega - \frac{1}{\mu_0 \sigma} \nabla^2\right) \boldsymbol{b_r} = \boldsymbol{v} \cdot \nabla B_{0r} + B_{0r} \frac{\partial \zeta}{\partial t}$$

- Amplitudes:
 - Eccentricity : 5x10⁻³ nT
 - Obliquity: 4 nT
 - Librational: 5 nT



 \rightarrow ~ 10% of response induced by Jovian B field induction



Librational tide magnetic perturbation at t = 0.25 (nT_{10})



Obliquity tide magnetic perturbation at t = 0.25 (nT) 50 0 50 -150 -100 -50 0 50 100 150 Longitude (1)

Outlook

- Predicted induced E- & B-field perturbations at Ganymede and Europa are detectable by JUICE instrumentation (RPWI & MAG)
- Our model allows us to detangle the contributions from Jovian B-field and ocean motion
- Will potentially allow us to constrain:
 - Ocean depth
 - Ocean conductance
 - Characteristics of internal heat sources
- Continuous measurements needed to constrain harmonics (order of ω_G & ω_E) higher harmonics of particular interest